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MAPPING VINYL CYANIDE AND OTHER NITRILES IN TITAN’S ATMOSPHERE USING ALMA

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ABSTRACT

Vinyl cyanide (C2H3CN) is theorized to form in Titan’s atmosphere through high-altitude photochemistry, and is of interest for astrobiology of cold planetary surfaces due to its predicted ability to form cell membrane-like structures (azotosomes) in liquid methane. In this work, we follow up on the initial spectroscopic detection of C2H3CN on Titan by Palmer et al. (2017) with the detection of three new C2H3CN rotational emission lines at sub-millimeter frequencies. These new, high-resolution detections have allowed for the first spatial distribution mapping of C2H3CN on Titan. We present simultaneous observations of C2H5CN, HC3N, and CH3CN emission, and obtain the first (tentative) detection of C3H8 (propane) at radio wavelengths. We present disk-averaged vertical abundance profiles, two-dimensional spatial maps, and latitudinal flux profiles for the observed gases. Similar to HC3N and C2H5CN, which are theorized to be short-lived in Titan’s atmosphere, C2H3CN is most abundant over the southern (winter) pole, whereas the longer-lived CH3CN is more concentrated in the north. This abundance pattern is consistent with the combined effects of high-altitude photochemical production, poleward advection, and the subsequent reversal of Titan’s atmospheric circulation system following the recent transition from northern to southern winter. We confirm that C2H3CN and C2H5CN are most abundant at altitudes above 200 km. Using a 300 km step model, the average abundance of C2H3CN is found to be 3.03 ± 0.29 ppb, with a C2H5CN/C2H3CN abundance ratio of 2.43 ± 0.26. Our HC3N and CH3CN spectra can be accurately modeled using abundance gradients above the tropopause, with fractional scale-heights of 2.05 ± 0.16 and 1.63 ± 0.02, respectively.

Keywords: planets and satellites: atmospheres — planets and satellites: individual (Titan) — techniques: imaging spectroscopy — techniques: interferometric

1. INTRODUCTION

Titan, Saturn’s largest moon, possesses the most substantial atmosphere of any known satellite in the Solar System. This atmosphere has been shown to contain a variety of organic compounds, including hydrocarbons as well as nitrogen- and oxygen-containing species (see the review by Bézard et al. 2014). Due to their often-large dipole moments and strong rotational transitions, nitriles (molecules containing a CN group) are readily detected in the millimeter/submillimeter range. By virtue of the rapid production in the upper atmosphere and relatively short chemical lifetimes of some nitriles (Wilson & Atreya 2004), these molecules can be used as powerful probes of the seasonally-variable chemistry and dynamics in Titan’s upper atmosphere (e.g. Teanby et al. 2012; Cordiner et al. 2014; Vinatier et al. 2015).

Vinyl cyanide, or acrylonitrile (C2H3CN), has come into focus due to the recent work of Stevenson et al. (2015), which showed, through theoretical liquid-phase calculations, that this molecule is one of the most favored to form thermodynamically stable membranes (azotosomes) in liquid methane at the surface temperature of Titan (approximately

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produce spectral line maps, the continuum was subtracted from the visibility amplitudes using the CASA uvcontsub task. The measured noise-level of the continuum-subtracted line-free data). The image pixel size was set to 0.1′′ and the spatial resolution (FWHM of the Gaussian restoring beam) was 0.78′′ on the sky. Information from the JPL Horizons system was used to transform the images from equatorial coordinates to projected linear distances with respect to the center of Titan, and to correct the -6.90 km s\(^{-1}\) Doppler shift with respect to the ALMA rest frame.

3. RESULTS

The observed spectrum is shown in Figure 1. This was produced by integrating over an elliptical region (with FWHM 2.3′′ × 1.5′′), the shape of which was determined by defining a pixel mask based on a flux threshold, which was varied until 90% of the total continuum flux was enclosed. The value of 90% was chosen in order to obtain the maximum signal from weak lines while excluding noisier data from the periphery of the image. The difference in line-to-continuum ratio for this region compared with a larger extraction aperture including 100% of the detected flux was found to be negligible for the purposes of our study. Spectral peaks were identified using the Splatalogue database for astronomical spectroscopy.

Routine interferometric observations of Titan were performed for the purpose of flux calibration as part of ALMA project 2013.1.00033.S on 2015 April 29. A single 3-minute integration of Titan was obtained at UT 09:19 using the band 7 receiver, covering frequencies in the range 335.3-349.2 GHz. Observations were made with thirty-nine antennae in the telescope array, which provided baselines in the range 15-349 m, with good u − v (Fourier) coverage of the sky. As with the previous work of Cordiner et al. (2014, 2015) and Palmer et al. (2017), even this short integration time was sufficient to reach a high enough sensitivity for the detection of weak molecular emission lines from Titan. The present study was further facilitated by the larger number of functioning antennas compared to the 20-30 available in previous studies (e.g. Cordiner et al. (2014, 2015) and Palmer et al. (2017)). The data used in our study were taken from a single spectral window in the upper receiver sideband, containing 3840 channels and with a channel spacing of 488 kHz which (after Hanning smoothing by the correlator) corresponds to a spectral resolution of 976 kHz. The low (0.57 mm) precipitable water vapor column and good phase stability provided excellent atmospheric conditions for these observations.

The data obtained from the ALMA Science Archive were processed in the NRAO CASA software (version 4.5.3) using the standard scripts provided by the Joint ALMA Observatory, as described by Cordiner et al. (2015). The measured continuum flux density for each baseline was scaled to match the Butler-JPL-Horizons 2012 Titan flux model. To produce spectral line maps, the continuum was subtracted from the visibility amplitudes using the CASA uvcontsub task. Imaging and deconvolution were performed using the clean task; the point-spread function was deconvolved using the Hogbom algorithm, with natural visibility weighting and a threshold flux level of 18 mJy (twice the RMS noise-level of the continuum-subtracted line-free data). The image pixel size was set to 0.1″ × 0.1″, and the spatial resolution (FWHM of the Gaussian restoring beam) was 1.22″ × 0.62″. For comparison, the angular diameter of Titan’s surface was 0.78″ on the sky. Information from the JPL Horizons system was used to transform the images from equatorial coordinates to projected linear distances with respect to the center of Titan, and to correct the -6.90 km s\(^{-1}\) Doppler shift with respect to the ALMA rest frame.

2 http://www.cv.nrao.edu/php/splat/
The spatial distribution of flux from each species is shown in the contour maps in Figure 3, in which the intensity of emission is indicated with an orange color scale and the black contours correspond to increments of $n\sigma$ (where $\sigma$ is the RMS noise in each map and $n$ is a constant factor). These maps were obtained by integrating over the widths of the lines in Table 1. For C$_2$H$_3$CN, only lines 1 and 3 were used in the map due to possible contamination of lines 2 and 4 with emission from C$_3$H$_8$ and CH$_3$CN, respectively. For CH$_3$CN, only line 3 was used, to probe CH$_3$CN from a single upper-state energy level (while excluding emission from the weaker, noisier lines of this species).

CH$_3$CN, and vibrationally excited HC$_3$N in the spectrum. In addition, two lines of C$_3$H$_8$ (propane) were tentatively detected, as well as two unidentified (U) lines at 347635 MHz and 347719 MHz.

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Figure 2. Close-up spectra of the three newly-detected $C_2H_3CN$ transitions (marked with asterisks).

Figure 3. Integrated emission contour maps for $C_2H_3CN$, $C_2H_5CN$, $HC_3N$, and $CH_3CN$. The coordinate scale is in Titan-projected distances and axes are aligned in the equatorial coordinate system. Titan’s surface is represented by the blue circle, with the white cross denoting the position of the phase center. Dashed white curves denote Titan’s equator and a blue dot-dashed line marks the orientation of the polar axis (2.6° counter-clockwise from vertical; north tilted toward the observer by 24.7°). The RMS noise level of each map is shown as hatched ellipses (Sun-target-observer) phase angle of 2.4°. The FWHM of the Gaussian restoring beam (1.42′′ x 0.62′′) and its orientation are shown as hatched ellipses.
Figure 4. Observed spectra with best-fitting NEMESIS models overlaid. Lower panels show the spectroscopic residuals (observation minus model).
Following the methodology of Cordiner et al. (2014, 2015), the observed spectrum was modeled using the line-by-line radiative transfer module of the NEMESIS atmospheric retrieval code (Irwin et al. 2008). For computational efficiency, the data were divided into three frequency ranges: (a) 347.2-348.0 GHz, (b) 348.2-348.6 GHz, and (c) 348.6-349.1 GHz – see Figure 4. The temperature profile was based on a combination of HASI (Fulchignoni et al. 2005) data and Cassini CIRS CH$_4$ limb-sounding data obtained during the T102, T105, T110, and T111 Titan flybys (between 2014 June and 2015 April). The CIRS temperatures as a function of latitude and altitude were retrieved following the same methodology as Acherberg et al. (2008). The disk-averaged CIRS temperature profile (covering altitudes $z \approx 60 – 550$ km) was combined with the HASI profile (covering $z = 0-150$ km) by feathering the profiles together in the range 60-100 km using a linearly-graded (weighted) average. The resulting temperature data are shown in Figure 5. As described in Section 4.2, latitudinal temperature variations may have a small impact on our retrieved abundances, but these are expected to be reduced to a negligible level as a result of the disk-averaging process. The abundances of nitrogen, methane, and carbon monoxide isotopologues were the same as those used by Teanby et al. (2013). To fit the $^{12}$CO line wing, which dominates the shape of the spectral slope in frequency range (a), the CO abundance was set to the more recent value of $5.06 \times 10^{-5}$ from Serigano et al. (2016). Propane (C$_3$H$_8$) has many transitions within our observed spectral range, but only two features (at 347.436 and 347.630 GHz) are strong enough to appear in the ALMA spectrum (and those only with a marginal significance); they were modeled using a constant C$_3$H$_8$ abundance of $4.2 \times 10^{-7}$ above the precipitation altitude as retrieved from CIRS observations by Nixon et al. 2009.

The model fluxes were calculated by integrating with a linear interpolation scheme over the radiances obtained across 35 impact parameters between the center of Titan’s disk and the top of the model atmosphere at 1000 km (see Teanby et al. 2013 for more details). The distribution of impact parameters was varied iteratively, starting with 10 and adding more until the improvement in retrieved abundances was less than 1/10 the abundance error. The sky-projected separation between individual rays was 500 km inside Titan’s disk (in the range 0-2575 km), 25 km in the range 2575-3075 km and 50 km in the range 3075-3575 km. The majority of the emission in our models originates in the altitude range 200-600 km; above 600 km, contributions are small due to exponential decay of the atmospheric density, but we extend up to 1000 km in order to avoid bias in our fractional scale height models (described below), which decrease with altitude at a slower rate. Each ray (impact parameter) in our model was weighted by a kernel defined by the shape of the flux extraction aperture (90% continuum flux contour) convolved with the instrumental point spread function. The continuum regions in the observed spectra (including the pseudo-continuum of the CO line wing) were scaled up to match the model fluxes to account for incomplete flux retrieval from the ALMA data, as well as correcting for any minor temperature differences between our model and that of the Butler-JPL-Horizons model used to calibrate the visibilities (see Section 2).

A step model (with a constant value above 300 km and zero below) was adopted for the C$_2$H$_3$CN and C$_2$H$_5$CN vertical abundance profiles (after Cordiner et al. 2015; Palmer et al. 2017), and fractional scale-height gradient models (with constant ratios of gas scale height to atmospheric scale height) were adopted for HC$_3$N and CH$_3$CN, with saturation laws taken from Loison et al. (2015). A Lorentzian broadening HWHM value of $\Gamma = 0.1$ cm$^{-1}$ bar$^{-1}$, and temperature exponent $\alpha = 0.75$ were used for HC$_3$N (taken from the HITRAN catalogue; Gordon et al. 2017). For C$_2$H$_4$CN, C$_2$H$_5$CN, and CH$_3$CN, standard values of $\Gamma = 0.075$ cm$^{-1}$ bar$^{-1}$, $\alpha = 0.5$ were used due to a relative lack of detailed laboratory measurements for these species. The abundances and fractional scale-height parameters $f_H$ were optimized by minimizing the sum of squares of the spectroscopic residuals, and good fits were obtained to the spectra for all species. The spectral fits are overlaid in Figure 4 for each of the three frequency ranges. The retrieved abundances at 300 km altitude are listed in Table 2, and the retrieved abundance profiles are plotted in Figure 6.

The adopted model abundance profiles were chosen to employ the least number of parameters required to obtain a reasonable fit to the observed spectra, given the noise. As shown by Cordiner et al. (2015) and Palmer et al. (2017), at a similar S/N, the C$_2$H$_3$CN and C$_2$H$_5$CN 300 km step models are not unique. Furthermore, above 300 km, the line wings become negligible. Thus, the observed spectra may also be fit using ‘gradient’ models or step-profiles with cutoffs at different altitudes. Based on the reduced $\chi^2$ values, which measure the goodness of fit, we find that models with a step height of >200 km provide a reasonable fit to our C$_2$H$_3$CN and C$_2$H$_5$CN observations; we adopt 300 km here to facilitate comparison with the results of Palmer et al. (2017). We note that our best-fitting ‘gradient’ model has an $f_H$ value (ratio of gas scale height to atmospheric scale height) of $1.8 \pm 0.2$, which is somewhat less than the value of $4.6 \pm 1.9$ derived by Palmer et al. (2017) and indicates a probable steepening of the C$_2$H$_3$CN...
Figure 5. Temperature data used to model the spectral lines. The North and South temperature profiles are averaged over a 0.25\" (Gaussian) beam centered over Titan’s northern and southern polar limbs.

Table 2. Retrieved molecular abundances at 300 km

<table>
<thead>
<tr>
<th>Species</th>
<th>Abundance (ppb)</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\text{C}_2\text{H}_3\text{CN}$</td>
<td>3.03 (0.29)</td>
</tr>
<tr>
<td>$\text{C}_2\text{H}_5\text{CN}$</td>
<td>7.37 (0.32)</td>
</tr>
<tr>
<td>$\text{HC}_3\text{N}^a$</td>
<td>4.50 (0.75)</td>
</tr>
<tr>
<td>$\text{CH}_3\text{CN}^a$</td>
<td>12.7 (0.75)</td>
</tr>
</tbody>
</table>

NOTE — Values in parentheses denote 1\(\sigma\) error margins. \(^a\)Fractional scale-heights \((f_H)\) for the best-fitting HC$_3$N and CH$_3$CN ‘gradient’ models are 2.05 ± 0.16 and 1.63 ± 0.02, respectively.

vertical abundance profile over the time period of about a year since those observations.

Vibrational emission from propane (C$_3$H$_8$) was first detected in Titan’s atmosphere by Voyager 1 (Maguire et al. 1981). Due to its theorized long photochemical lifetime (~ 300 years; Wilson & Atreya 2004), propane may be expected to be quite well mixed throughout Titan’s atmosphere. It shows relatively little spatial variation compared with other species (e.g. Vinatier et al. 2015), so we approximate the C$_3$H$_8$ distribution in our radiative transfer model using a vertical profile with a constant abundance above the saturation altitude of 60 km. **Pressure broadening coefficients of $\Gamma = 0.12 \text{ cm}^{-1} \text{ bar}^{-1}$, $\alpha = 0.5$ were adopted based on the typical values for rovibrational transitions of C$_3$H$_8$ (see Gordon et al. 2017).** Allowing the C$_3$H$_8$ abundance to vary as a free parameter in our model fits, we obtained a best-fitting C$_3$H$_8$ mixing ratio of $(8.5 \pm 0.6) \times 10^{-7}$ based on the two strongest rotational lines of this molecule in our observed spectral region. Although this is about a factor of two greater than the value of $(4.2 \pm 0.5) \times 10^{-7}$ reported by Nixon et al. (2009) using Cassini CIRS observations, our fitted value should be treated with extreme caution due to the fact that the two propane lines are both blended with other spectral features, leading
Figure 6. Vertical abundance profiles derived from radiative transfer modeling, with 1σ error envelopes.

Table 3. Parameters of Gaussian fits to spatial profiles

<table>
<thead>
<tr>
<th>Species</th>
<th>S pos. (km)</th>
<th>N pos. (km)</th>
<th>Flux Ratio (South/North)</th>
</tr>
</thead>
<tbody>
<tr>
<td>C_2H_3CN</td>
<td>-2569 (302)</td>
<td>2842 (770)</td>
<td>2.83 (0.77)</td>
</tr>
<tr>
<td>C_2H_5CN</td>
<td>-2243 (115)</td>
<td>2144 (191)</td>
<td>1.52 (0.10)</td>
</tr>
<tr>
<td>HC_3N</td>
<td>-2790 (94)</td>
<td>1629 (292)</td>
<td>3.07 (0.29)</td>
</tr>
<tr>
<td>CH_3CN</td>
<td>-1930 (98)</td>
<td>2096 (60)</td>
<td>0.67 (0.02)</td>
</tr>
</tbody>
</table>

NOTE — Distances represent the Gaussian peak positions along Titan’s sky-projected polar axis (from south to north), with the subobserver point as the origin. Values in parentheses denote 1σ error margins. Gaussian FWHM were fixed at the instrumental resolution of 4600 km.

to uncertainty in their observed strengths and profiles. The spectral regions surrounding the transitions are shown in Figure 7; the 23_0,23 − 22_1,22 line lies on the high-frequency side of a C_2H_3CN line, and the 23_1,23 − 22_0,22 line appears to be blended with the wing of an unidentified spectral feature at 347.635 GHz. In Figure 7, we overlay (1) the results of our best-fitting C_3H_8 radiative transfer model, (2) the predicted spectrum using the abundance obtained by Nixon et al. (2009), and (3) a model with no C_3H_8. We find that models including C_3H_8 provide a good fit to the observations of the 23_0,23 − 22_1,22 line whereas models without any C_3H_8 do not, as evidenced by better reduced χ² values (1.09 with the abundance from Nixon et al. (2009) and 1.06 with the best fitting abundance, versus 1.73 without), which confirms the likely presence of this species in our data. However, due to the line-blending issues, we are unable to place firm constraints on the C_3H_8 abundance or its vertical profile. It should also be noted that comparison of our retrieved abundance to that of Nixon et al. (2009) may be complicated by the fact that the value reported by Nixon et al. (2009) represents an average over latitudes −30° to 30° and does not include the polar regions. The CIRS
Figure 7. Close-up view of the spectral regions surrounding the two strongest $\text{C}_3\text{H}_8$ transitions in our data. Three different model curves are overlaid (see text), showing that $\text{C}_3\text{H}_8$ is likely to be present in our data, albeit blended with other transitions, which preclude a robust abundance estimate for this molecule. In particular, for the feature at 347.436 GHz, the probabilities that Gaussian (thermal) noise is responsible for the observed discrepancy between the model and data are as follows: No $\text{C}_3\text{H}_8$ model — 0.30; Nixon et al. (2009) $\text{C}_3\text{H}_8$ model — 0.13; Best-fitting $\text{C}_3\text{H}_8$ model — $1.8 \times 10^{-5}$. Residuals (observation minus model) are also plotted for the $2\nu_{0,23} - 2\nu_{1,22}$ line for the model with no $\text{C}_3\text{H}_8$, and the best-fitting model. Unfortunately, blending with the unidentified feature at 347.635 GHz precludes a similar statistical analysis of the feature at 347.630 GHz. However, the size of the feature predicted by the model is consistent with the observations.
and ALMA C$_2$H$_8$ observations are also likely to be sensitive to emission from different altitudes, due to the greatly differing spectral resolutions of these instruments.

4. DISCUSSION

4.1. Vertical Mixing Ratio Profiles

The detected C$_2$H$_3$CN transitions are optically thin and their strengths exhibit a weak temperature dependence in the range ($\approx$ 160 – 180 K) found in Titan’s atmosphere above 300 km. The vertical abundance profile for this species therefore implies that C$_2$H$_3$CN is present with an average abundance of 3.03 ± 0.29 ppb in Titan’s atmosphere above 300 km. This value is in agreement, within error, with the abundance of 2.83 ± 0.24 ppb derived by Palmer et al. (2017). Both observations confirm substantial production of C$_2$H$_3$CN in Titan’s upper/middle atmosphere, as predicted by the model of Loison et al. 2015. Given the downward mixing of photochemical products over time, the relatively low tropospheric C$_2$H$_3$CN abundance implies a short chemical lifetime for this molecule compared to those of more stable species such as C$_2$H$_2$, C$_2$H$_6$, HCN, and CH$_3$CN, which remain abundant as they travel into the lower stratosphere.

Our retrieved abundance for C$_2$H$_5$CN using a 300 km step model (7.37 ± 0.32 ppb) is close to the values obtained in the previous work of Cordiner et al. (2015) and Palmer et al. (2017) (9.25 ppb and 7.2 ppb, respectively). The C$_2$H$_5$CN/C$_2$H$_3$CN abundance ratio of 2.43 ± 0.26 is also consistent with the value of 2.5 ± 0.3 found by Palmer et al. (2017). Overall, the latest chemical models (see Loison et al. (2015) and Dobrijevic et al. (2016)), do not fit the C$_2$H$_3$CN results well, but provide a better fit for C$_2$H$_5$CN. Our newly determined C$_2$H$_5$CN abundance confirms a tendency for these models to over-estimate the stratospheric C$_2$H$_5$CN abundance. This could support the hypothesis that additional (or more rapid) destruction reactions need to be added to chemical networks. In particular, Dobrijevic et al. (2016) suggested that although the inclusion of ion chemistry reduced their over-estimation of the C$_2$H$_5$CN abundance, the remaining over-estimation is suggestive of a destruction mechanism which still remains unaccounted for. Alternatively, the discrepancy could be due to the fact that no current photochemical models have incorporated 2D dynamics yet. Further modeling will be required to elucidate whether and which of these changes could lead to improved modeled abundances for C$_2$H$_5$CN in Titan’s atmosphere.

Our disk-averaged ALMA abundance profile for CH$_3$CN can be compared with that obtained by Marten et al. (2002),
based on observations using the IRAM 30-m telescope. Our abundance profiles are generally in good agreement, except in the lower regions of the atmosphere. Below 200 km, the abundance profile of Marten et al. (2002) drops off much more rapidly with altitude (see their Figure 12). This could be a manifestation of stronger downward mixing of photochemical products around the time of our observations, which could be explainable as a result of the closer temporal proximity to Titan’s 2017 solstice, resulting in stronger atmospheric downwelling over the winter pole. Comparing our results to the models of Loison et al. (2015) and Dobrijevic et al. (2016), the modeled CH$_3$CN abundances increase more rapidly with altitude than the values we retrieve. As in the case of C$_2$H$_3$CN, this discrepancy may indicate that a destruction mechanism remains unaccounted for in the lower atmosphere. Dobrijevic et al. (2016) suggest that these processes may include meridional transport or sticking to aerosol grains. There are also significant differences between our retrieved CH$_3$CN profile and the model of Krasnopolsky (2009), with an observed abundance at 500 km of 6 $\times$ 10$^{-8}$ compared to the modeled value of 1 $\times$ 10$^{-6}$ (see his Figure 9). This suggests that the chemical production and destruction scheme for this molecule on Titan still remains to be fully understood.

In comparison to previously reported results for HC$_3$N (Marten et al. 2002; Teanby et al. 2007; Vinatier et al. 2010; Cordiner et al. 2014), our retrieved abundance for this molecule is greater in the lower stratosphere (at altitudes up to about 300 km). Above this level, our retrieved profile matches previous profiles well, with the exception of that retrieved by Cordiner et al. (2014), who obtained a profile with a lower abundance of HC$_3$N above this altitude. When compared to models by Loison et al. (2015), our retrieved profile falls within their 90% uncertainty interval. Given this good agreement with modeling results, it would appear that our results support the proposed formation of HC$_3$N via the reaction of CN with C$_2$H$_2$. Nonetheless, Dobrijevic et al. (2016) suggest that this molecule’s abundance may also be affected by the aforementioned missing destruction mechanisms; if this were accounted for, modeled abundances in the lower troposphere may in fact match our retrieved abundances with an even better fit.

4.2. Molecular Emission Maps

The additional C$_2$H$_3$CN emission lines detected not only increase our confidence in the previous detection by Palmer et al. (2017), but have allowed the first mapping of the spatial distribution of this molecule. Our C$_2$H$_3$CN map (Figure 3) shows a flux peak at 5$\sigma$ confidence in Titan’s southern hemisphere, consistent with enrichment over the south polar region. A similar south-polar enrichment is also evident for HC$_3$N and C$_2$H$_5$CN. Polar enrichments have previously been observed for a number of hydrocarbons and nitriles by Cassini and ALMA (Teanby et al. 2013; Cordiner et al. 2015; Vinatier et al. 2015). It has been theorized that as Titan transitions from northern winter in 2004 towards the southern winter solstice in May 2017, the main atmospheric circulation cell, responsible for the redistribution of photochemical products from mid-latitudes towards the poles, should reverse its direction (e.g. Teanby et al. 2012). The abundances of these molecules are thus expected to begin increasing in the southern polar region, while molecules concentrated in the north polar region would begin to undergo photochemical destruction, with a corresponding decay in the previous northern peak at a rate in accordance with the photochemical lifetime and diffusion rate of each molecule.

A clearer view of the latitudinal variability of the observed gases is shown in Figure 8, where we present the peak-normalized emission line fluxes as a function of sky-projected distance along Titan’s polar axis from south to north. Two-component Gaussian fits with freely-variable positions and intensities for the Gaussian components have been performed to these profiles, but with FWHM fixed at the instrumental resolution of 4600 km. These fits are overlaid in Figure 8 with dashed lines, and the positions and relative strengths of the components are given in Table 3. For all species, a reasonably good fit was obtained within the noise. The positions of the fit components relative to Titan’s 2575 km radius are consistent with the majority of the emission originating from near the poles, and from within a region much smaller than the spatial resolution element. Such confinement, whereby the majority of the gas is located within 20° of the pole, has previously been observed by Cassini for HC$_3$N and other short-lived species (e.g. Teanby et al. 2008). Discrepancies between the fits and observations may be due to noise, or to the presence of additional flux components.

The one-dimensional spatial flux profiles highlight the pronounced southern peak for HC$_3$N (with a contrast factor of about three relative to the northern peak). The factor of ~3 southern enhancement of the C$_2$H$_3$CN abundance appears to be similar to that of HC$_3$N, whereas the somewhat smaller factor of ~1.5 enhancement in C$_2$H$_5$CN implies either a slower production rate for C$_2$H$_3$CN in the south, or a slower destruction rate in the north compared with HC$_3$N and C$_2$H$_5$CN. This result is in contrast to the ALMA observations from 2012 by Cordiner et al. (2015), which showed C$_2$H$_5$CN to be enriched in the south while HC$_3$N was stronger in the north, leading to the conclusion that C$_2$H$_5$CN has the shorter chemical lifetime. Only three terrestrial years later, C$_2$H$_3$CN remains concentrated in the south, while the location of the HC$_3$N peak appears to have shifted from the north to the south. However, Cordiner...
et al. (2015) did not present information on the vertical distribution of HC$_3$N, so it is possible that the northern HC$_3$N peak in that study is caused by the presence of colder, higher-pressure gas, at lower altitudes than that probed by the vibrationally-excited HC$_3$N lines in our study. Continued spectral, spatial, and temporal monitoring of these molecules will be required to further elucidate the details of their relative production, destruction, and transport rates as a function of altitude.

Significant stratospheric temperature differences have previously been identified between Titan’s northern and southern polar regions, which could cause differences in the emission line strengths as a result of (1) the temperature dependence of the molecular excitation, (2) the Planck radiation function, and (3) the molecular partition functions. To assess the possible impact of these temperature effects, we extracted spatially-resolved CH$_4$ temperature data for the complete latitudinal range from the T102, T105, T110, and T111 Cassini flybys (between 2014 June and 2015 April), as explained in Section 3. These temperature data were averaged within 0.25" Gaussian beams centered over the north and south polar limbs, chosen to cover the approximate extent of the polar latitudinal regions within which our observed molecules are likely to be concentrated (see Teanby et al. 2008). Between $z = 250$-400 km, the temperature difference between the north and south limbs was found to be 5-8 K (see Figure 5), and at all other altitudes the temperature difference was less than 5 K. The resulting temperature-dependence of our detected emission line strengths was found to be $< 20\%$ for HC$_3$N, $< 10\%$ for CH$_3$CN, and negligible for the other species. Combined with the fact that the observed emission is quite optically thin, the structure in the spatial flux profiles can thus be interpreted as primarily due to latitudinal variations in the observed gas abundances.

In contrast to the other species observed in the present study, the CH$_3$CN flux peaks in Titan’s northern hemisphere, consistent with a greater concentration of this species over the north pole. Given the near-complete transition to northern summer following the northern spring equinox in 2009, this remnant gas concentration indicates that CH$_3$CN is more resistant to photochemical destruction than the other species observed, which agrees with the relatively long ($\approx 8$ yr) theorized photochemical lifetime for CH$_3$CN compared to the shorter ($\lesssim 1$ yr) lifetimes for C$_2$H$_5$CN, C$_2$H$_5$N, and HC$_3$N (see e.g. Krasnopolsky 2009). Our latitudinal profiles are qualitatively consistent with the model results of (Loison et al. 2015), who determined longer lifetimes for CH$_3$CN and C$_2$H$_5$CN than for HC$_3$N and C$_2$H$_5$N. The differences in these molecules’ lifetimes in Titan’s atmosphere may be explained by their degree of saturation: CH$_3$CN and C$_2$H$_5$CN are fully saturated in the sense that all carbon atoms are bonded to four other atoms (with the exception of the carbon in the nitrile group), whereas HC$_3$N and C$_2$H$_5$N are unsaturated, containing carbon atoms that are double- or triple-bonded. Saturated molecules are less reactive, which explains the longer lifetimes of CH$_3$CN and C$_2$H$_5$CN relative to HC$_3$N and C$_2$H$_5$N. Furthermore, CH$_3$CN and C$_2$H$_5$CN are less rapidly photodissociated than HC$_3$N and C$_2$H$_5$CN due to line shielding by CH$_4$ (J. C. Loison 2017, private communication). A similar relationship in chemical lifetimes between HC$_3$N, C$_2$H$_5$N, and CH$_3$CN is produced in the latest chemical model of V. Vuitton (2017, private communication). This model also points to photolysis rates as an important factor in the chemical lifetimes for these species. In particular, HC$_3$N is more efficiently photolysed than CH$_3$CN by the longer wavelength photons found in the stratosphere.

The importance of Titan’s polar tilt with respect to the line of sight (currently 24.7°) needs to be considered when interpreting latitudinal flux profiles. Whereas Titan’s north polar surface was in full view of Earth at the epoch of our observations, the region south of $-65^\circ$ was obscured behind Titan’s disk. The detailed distribution of Titan’s CH$_3$CN is currently not well-known, but if a significant proportion is concentrated at low altitudes over the pole, this could be responsible for some of the apparent northerly enhancement. Conversely, the southerly enhancement for the other observed species is likely to be even more pronounced if the south pole were in full view.

5. CONCLUSION

We have presented here observations of three new transitions of vinyl cyanide, C$_2$H$_4$CN, strengthening our confidence in the previously reported spectroscopic detection of this molecule by Palmer et al. (2017). We also report a tentative spectroscopic detection of propane, C$_3$H$_8$.

Abundances were derived for C$_2$H$_5$CN, C$_2$H$_5$N, HC$_3$N, and CH$_3$CN using radiative transfer modeling. These results show generally a good agreement with prior observations. The retrieved abundance for C$_2$H$_5$CN is $3.03 \pm 0.29$ ppb above 300 km, and our models support previous indications that this molecule is produced efficiently on Titan at high altitudes. Discrepancies between our retrieved abundance profiles and chemical models for CH$_3$CN, C$_2$H$_5$CN and C$_2$H$_5$N highlight the need for improved understanding of the chemistry of these species in planetary atmospheres.

We were able to produce the first spatially resolved map of the distribution of C$_2$H$_5$CN on Titan, as well as new maps of C$_2$H$_5$CN, HC$_3$N, and CH$_3$CN for the 2015 April epoch. Comparing these to maps produced in prior epochs, we find support for a global redistribution of atmospheric nitriles, in accordance with the hypothesized seasonally-driven
global circulation cell direction change. Furthermore, the observed changes lend support to chemical model predictions that HC$_3$N and C$_2$H$_3$CN have a shorter lifespan in Titan’s atmosphere than C$_2$H$_5$CN and CH$_3$CN. The continued northern enrichment of CH$_3$CN despite the change in season supports a relatively longer chemical lifespan than for the other observed species.

Following our detection and mapping of C$_2$H$_3$CN and other nitriles using ALMA, additional detailed mapping (and chemical modeling) of the time-variability of these species is warranted to further understand their photochemistry in primitive planetary atmospheres.

This work was supported by The Goddard Center for Astrobiology and by the National Science Foundation under Grant No. AST-1616306. It makes use of ALMA data set ADS/JAO.ALMA#2013.1.00033.S. ALMA is a partnership of ESO (representing its member states), NSF (USA), and NINS (Japan), together with NRC (Canada), and NSC and ASIAA (Taiwan), in cooperation with the Republic of Chile. The Joint ALMA Observatory is operated by ESO, AUI/NRAO, and NAOJ. The National Radio Astronomy Observatory is a facility of the National Science Foundation operated under cooperative agreement by Associated Universities, Inc. CAN and SBC were supported by the NASA HQ Science Innovation Fund for their portion of the work in this paper.

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